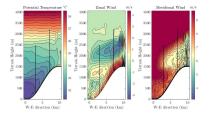
# A new $K - \varepsilon$ turbulence parameterization for mesoscale meteorological models

# <u>A. Zonato<sup>1</sup></u>, A. Martilli<sup>2</sup>, P. A. Jimenez<sup>3</sup>, J. Dudhia<sup>3</sup>, D. Zardi<sup>1</sup> & L. Giovannini<sup>1</sup>

<sup>1</sup>University of Trento, Trento, Italy <sup>2</sup>CIEMAT, Madrid, Spain <sup>3</sup>NCAR, Boulder, Colorado



Denver, January 12, 2023

Improve the **reproduction** of boundary layer dynamics at the **mesoscale**, through **novel turbulence closures** 

Improve the **reproduction** of boundary layer dynamics at the **mesoscale**, through **novel turbulence closures** 

#### Why?

Currently mesoscale models adopt 1.5 order  $K - \ell$  1D turbulence closures, mostly for numerical stability reasons.

Improve the **reproduction** of boundary layer dynamics at the **mesoscale**, through **novel turbulence closures** 

#### Why?

Currently mesoscale models adopt 1.5 order  $K-\ell$  1D turbulence closures, mostly for numerical stability reasons.

#### Problems...

× Correctly define the length scale for TKE and dissipation, especially in <u>complex</u> and <u>heterogeneous</u> terrain, where the "memory effect" can be relevant

Improve the **reproduction** of boundary layer dynamics at the **mesoscale**, through **novel turbulence closures** 

#### Why?

Currently mesoscale models adopt 1.5 order  $K-\ell$  1D turbulence closures, mostly for numerical stability reasons.

#### Problems...

- × Correctly define the length scale for TKE and dissipation, especially in complex and heterogeneous terrain, where the "memory effect" can be relevant
- $\times~\ell$  is commonly obtained from measurements/LES in **flat** terrain

Improve the **reproduction** of boundary layer dynamics at the **mesoscale**, through **novel turbulence closures** 

#### Why?

Currently mesoscale models adopt 1.5 order  $K-\ell$  1D turbulence closures, mostly for numerical stability reasons.

#### Problems...

- × Correctly define the length scale for TKE and dissipation, especially in complex and heterogeneous terrain, where the "memory effect" can be relevant
- $\times~\ell$  is commonly obtained from measurements/LES in **flat** terrain

#### Idea!

Employing a  $K-\varepsilon$  closure in order to avoid to define a mixing length scale  $$^{1/12}$$ 

PBL equations

# The $K - \varepsilon$ turbulence closure (1.5 order)

Mixing coefficient





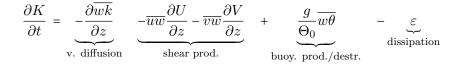
#### The $K - \varepsilon$ turbulence closure (1.5 order)

Mixing coefficient





© PROGNOSTIC Turbulent Kinetic Energy (K) equation



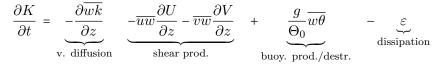
#### The $K - \varepsilon$ turbulence closure (1.5 order)

Mixing coefficient



 $\nu_{\chi} = \mathbf{c}_{\mu} \frac{\mathbf{K}^2}{2}$ 

© PROGNOSTIC Turbulent Kinetic Energy (K) equation



© **PROGNOSTIC** Dissipation Rate ( $\varepsilon$ ) equation

$$\frac{\partial \varepsilon}{\partial t} = \underbrace{-\frac{1}{\sigma_{\varepsilon}} \frac{\partial \overline{\varepsilon} \overline{w}}{\partial z}}_{\text{v. diffusion}} - \underbrace{\left[ \frac{c_1 \left( \overline{u} \overline{w} \frac{\partial U}{\partial z} + \overline{v} \overline{w} \frac{\partial V}{\partial z} \right) - \frac{c_3 g}{\Theta_0} \overline{w} \overline{\theta} \right]}_{\text{shear+buoy. prod./destr.}} \underbrace{\frac{\varepsilon}{K}}_{1/\tau} - \underbrace{\frac{c_2 \varepsilon}{K}}_{1/\tau}$$

#### Tuning the standard $K-\varepsilon$ closure

a) Prandtl number for the mixing coefficient (Hong et al., 2006)

$$\nu_H = \frac{\nu_M}{Pr}$$
  $Pr = 1 + (Pr_0 - 1) \exp\left[\frac{-3(z - 0.1h)^2}{h^2}\right]$ 

#### Tuning the standard $K - \varepsilon$ closure

a) Prandtl number for the mixing coefficient (Hong et al., 2006)

$$\nu_H = \frac{\nu_M}{Pr}$$
  $Pr = 1 + (Pr_0 - 1) \exp\left[\frac{-3(z - 0.1h)^2}{h^2}\right]$ 

b) Dissipation dependence on the eddy scale (Zeng et al., 2020)

Buoy prod = Buoy prod + 
$$c_4 \min\left(1, \sqrt{\frac{Ri}{c_5}}\right) N\varepsilon$$

#### Tuning the standard $K - \varepsilon$ closure

a) Prandtl number for the mixing coefficient (Hong et al., 2006)

$$\nu_H = \frac{\nu_M}{Pr}$$
  $Pr = 1 + (Pr_0 - 1) \exp\left[\frac{-3(z - 0.1h)^2}{h^2}\right]$ 

b) Dissipation dependence on the eddy scale (Zeng et al., 2020)

Buoy prod = Buoy prod + 
$$c_4 \min\left(1, \sqrt{\frac{Ri}{c_5}}\right) N\varepsilon$$

c) Counter-gradient term for the heat flux

1) 
$$\overline{w\theta} = -\nu_H \left(\frac{\partial\Theta}{\partial z} - \gamma\right) \qquad \gamma = C \frac{\overline{w\theta}_s}{w_\star h} \qquad (NL)$$

#### Tuning the standard $K - \varepsilon$ closure

a) Prandtl number for the mixing coefficient (Hong et al., 2006)

$$\nu_H = \frac{\nu_M}{Pr}$$
  $Pr = 1 + (Pr_0 - 1) \exp\left[\frac{-3(z - 0.1h)^2}{h^2}\right]$ 

b) Dissipation dependence on the eddy scale (Zeng et al., 2020)

Buoy prod = Buoy prod + 
$$c_4 \min\left(1, \sqrt{\frac{Ri}{c_5}}\right) N\varepsilon$$

c) Counter-gradient term for the heat flux

1) 
$$\overline{w\theta} = -\nu_H \left(\frac{\partial\Theta}{\partial z} - \gamma\right) \qquad \gamma = C \frac{\overline{w\theta}_s}{w_\star h}$$
 (NL)

2) 
$$\overline{w\theta} = -\nu_H \frac{\partial \Theta}{\partial z} + \Phi_{cg}(K_\theta) \quad \frac{\partial K_\theta}{\partial t} = -\frac{\partial \overline{wK_\theta}}{\partial z} - \overline{w\theta} \frac{\partial \Theta}{\partial z} - \varepsilon_\theta \quad (\mathbf{L})$$

# Numerically speaking...

# Numerically speaking...

Neglecting vertical diffusion + variable change

$$X = \frac{K}{\varepsilon} \qquad Y = \varepsilon^{\alpha} K^{\beta} \qquad A = c_{\mu} \left( S^2 - \frac{N^2}{Pr} \right) \qquad B = c_{\mu} \left( c_1 S^2 - c_3 \frac{N^2}{Pr} \right)$$

#### Numerically speaking...

Neglecting vertical diffusion + variable change

$$X = \frac{K}{\varepsilon} \qquad Y = \varepsilon^{\alpha} K^{\beta} \qquad A = c_{\mu} \left( S^2 - \frac{N^2}{Pr} \right) \qquad B = c_{\mu} \left( c_1 S^2 - c_3 \frac{N^2}{Pr} \right)$$

$$\frac{\partial X}{\partial t} = -CX^2 + (c_2 - 1) \tag{1a}$$

$$\frac{\partial \ln Y}{\partial t} = (\alpha A + \beta B) X - (\alpha + \beta c_2) \frac{1}{X}$$
(1b)

...analytical solution!!

#### Numerically speaking...

Neglecting vertical diffusion + variable change

$$X = \frac{K}{\varepsilon} \qquad Y = \varepsilon^{\alpha} K^{\beta} \qquad A = c_{\mu} \left( S^2 - \frac{N^2}{Pr} \right) \qquad B = c_{\mu} \left( c_1 S^2 - c_3 \frac{N^2}{Pr} \right)$$

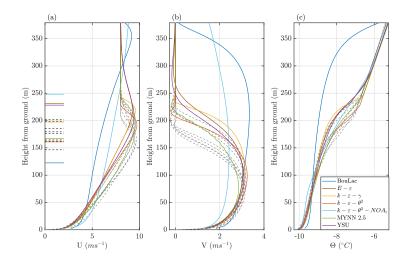
$$\frac{\partial X}{\partial t} = -CX^2 + (c_2 - 1) \tag{1a}$$

$$\frac{\partial \ln Y}{\partial t} = (\alpha A + \beta B) X - (\alpha + \beta c_2) \frac{1}{X}$$
(1b)

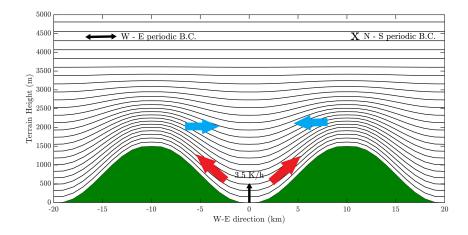
#### ...analytical solution!!

After solving the system, **diffusion** is calculated, and **advection** is applied to  $\mathbf{K}$ ,  $\varepsilon$  and  $\theta^2$ 

#### 1D - Stable - GABLS2 - LES (2 m) vs RANS (1 km)

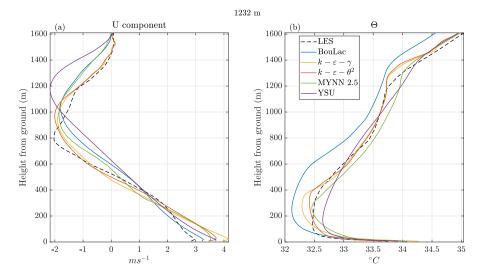


#### 2D - LES and RANS simulation Set-Up

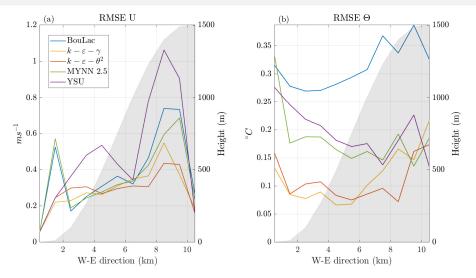


Setup

# 2D - Complex terrain - LES (50 m) vs RANS (1 km)

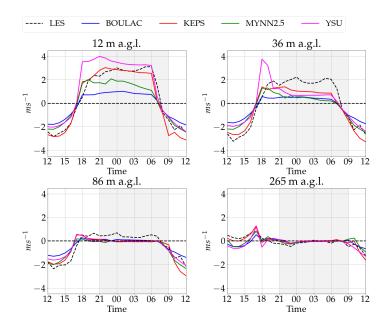


# 2D - Complex terrain - LES (50 m) vs RANS (1 km)



Results

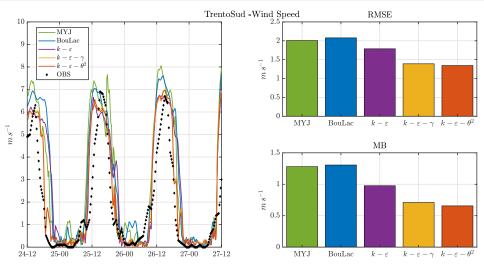
# 3D - Complex terrain - LES (100 m) vs RANS (1 km)



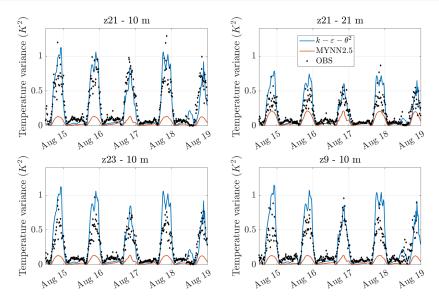
Results

9/12

#### Real 3D - SIM vs OBS - 10-m wind speed



#### **Real 3D** - WFIP2 - Temperature Variance



Results

✓ The new tuned  $k - \varepsilon$  outperforms other PBL models. The best results are shown by the scheme with temperature variance prognostic equation. We can deduce that:

- ✓ The new tuned  $k \varepsilon$  outperforms other PBL models. The best results are shown by the scheme with temperature variance prognostic equation. We can deduce that:
  - a) The inclusion of a prognostic equation for  $\varepsilon$  ("memory effect") improves the results with respect to the use of a diagnostic  $\ell$ ;
  - b) Model **locality** improves the simulations, especially at increasing levels of complexity;

- ✓ The new tuned  $k \varepsilon$  outperforms other PBL models. The best results are shown by the scheme with temperature variance prognostic equation. We can deduce that:
  - a) The inclusion of a prognostic equation for  $\varepsilon$  ("memory effect") improves the results with respect to the use of a diagnostic  $\ell$ ;
  - b) Model **locality** improves the simulations, especially at increasing levels of complexity;
- $\checkmark$  (Hopefully) available in the next **WRF** version!

- ✓ The new tuned  $k \varepsilon$  outperforms other PBL models. The best results are shown by the scheme with temperature variance prognostic equation. We can deduce that:
  - a) The inclusion of a prognostic equation for  $\varepsilon$  ("memory effect") improves the results with respect to the use of a diagnostic  $\ell$ ;
  - b) Model **locality** improves the simulations, especially at increasing levels of complexity;
- $\checkmark$  (Hopefully) available in the next **WRF** version!

- ✓ The new tuned  $k \varepsilon$  outperforms other PBL models. The best results are shown by the scheme with temperature variance prognostic equation. We can deduce that:
  - a) The inclusion of a prognostic equation for  $\varepsilon$  ("memory effect") improves the results with respect to the use of a diagnostic  $\ell$ ;
  - b) Model **locality** improves the simulations, especially at increasing levels of complexity;
- $\checkmark$  (Hopefully) available in the next **WRF** version!

Reference: A. Zonato, A. Martilli, P. A. Jimenez, J. Dudhia, D. Zardi & L. Giovannini, A new  $K - \varepsilon$  turbulence parameterization for mesoscale meteorological models, Monthly Weather Review, 2022.

Acknowledgments and funding: Atmospheric boundary-layer modeling over complex terrain (ASTER) project.